

Empirical methods for the assessment of seismic system sensitivity

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Abstract

One of the basic performance characteristics of the seismic system is its sensitivity. An appreciation of the spatial variation of m_{min} is important for the correct interpretation of seismic data and the planning of future expansion of the seismic system. This paper will discuss the empirical evaluation of the system sensitivity applied to the data from two Australian mines. The method discussed here is reliant on enough seismic data being available and is per se not applicable to greenfields sites. A gridless method for the evaluation of m_{min} with respect to the distance to sensors is proposed as a stable and relatively easy first order approach to the evaluation of the system sensitivity that overcomes some of the problems of the other methods. This method is easy to understand and facilitate the conceptualisation of the influence of the sensor locations on the system sensitivity.

1 Introduction

One of the basic characteristics of the seismic system is its sensitivity. The sensitivity of the seismic system is defined in this paper as the minimum magnitude event that is reliably recorded by the system, and is designated by m_{min} . This is also referred to as the magnitude of completeness of the record (Sagar and Leonard 2007; Wiemer and Wyss, 2000), as this is the magnitude above which all events are recorded and therefore, a complete dataset exists.

An appreciation of the spatial variation of m_{min} is important for the correct interpretation of seismic data. Planning of future expansion of the seismic system also requires an appreciation of the current achieved sensitivity, as this allows the planning of expansions to the system to be focused on the areas of interest.

Figure 1 illustrates the effect of the spatial variation of m_{min} over the mine. The figure shows what is generally referred to as the magnitude–time history plot (Hudyma et al., 2003). The scatter plot shows the magnitude of events over time, while the line shows the cumulative number of events with time. For the first period up to August 2006, the smallest events recorded were about $M_L=-2$ with a m_{min} of about -1.5. After this period there is a sudden increase in system sensitivity (drop in the smallest recorded event size) due to the fact that more sensors were installed to better cover the area of highest seismic activity. Another sudden increase in the system sensitivity occurs around November 2007 when a few more sensors were added to the deeper levels. Between August 2006 and November 2007, there is a gradual reduction in the system sensitivity. This apparent reduction in the system sensitivity is not due to any change in the system configuration over time, but due to the mining progressively migrating deeper and ‘out of’ the array into areas not well covered by the array. This is illustrated in Figure 2.

This paper discusses the empirical evaluation of the system sensitivity applied to the data of two Australian mines. The method discussed here is reliant on enough seismic data being available and is *per se* not applicable to greenfields sites. For greenfields sites, the evaluation of the future system sensitivity will have to be based on theoretical methods only.

Although the method discussed here, being empirical, is only applicable to the dataset, associated volume and seismic system configuration at the time, it can be used for the first order design of system expansion.

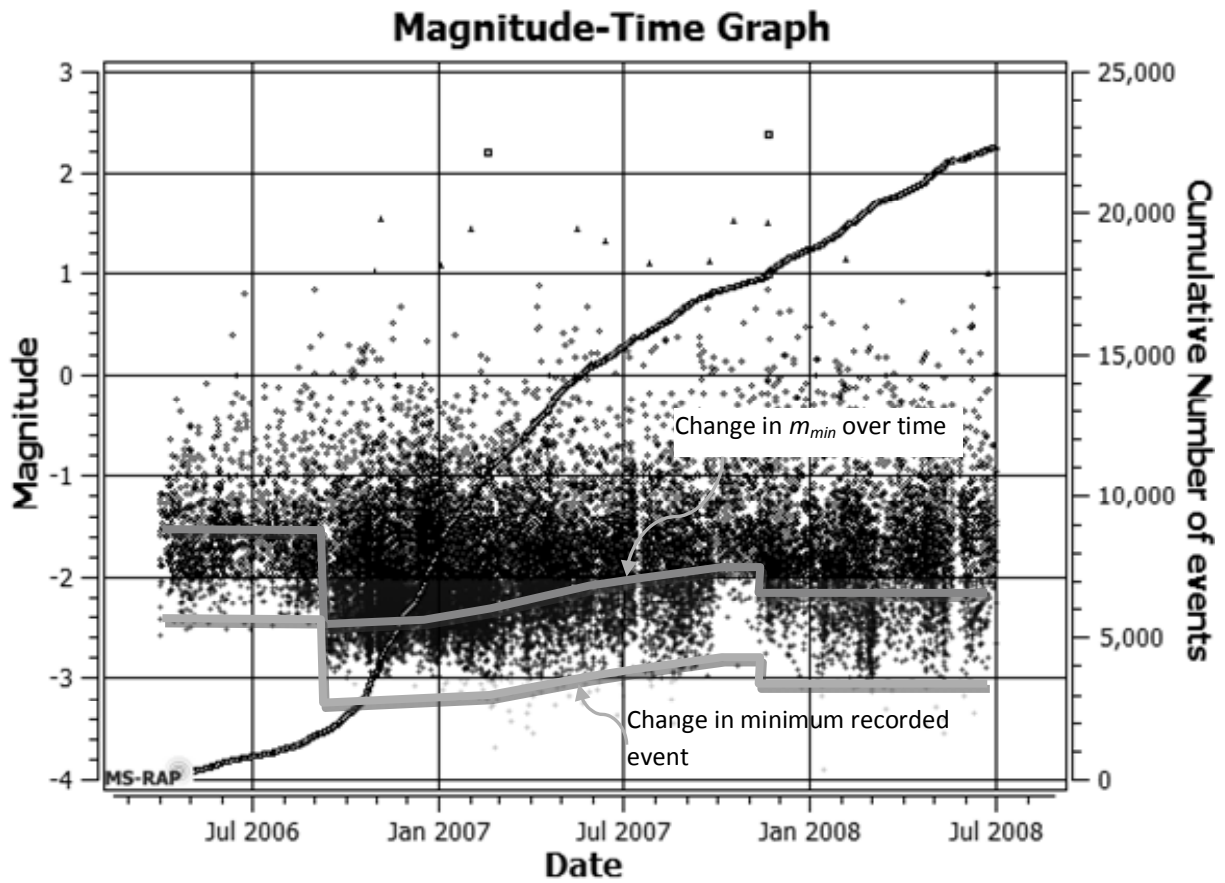


Figure 1 Magnitude-time graph for data from an Australian mine

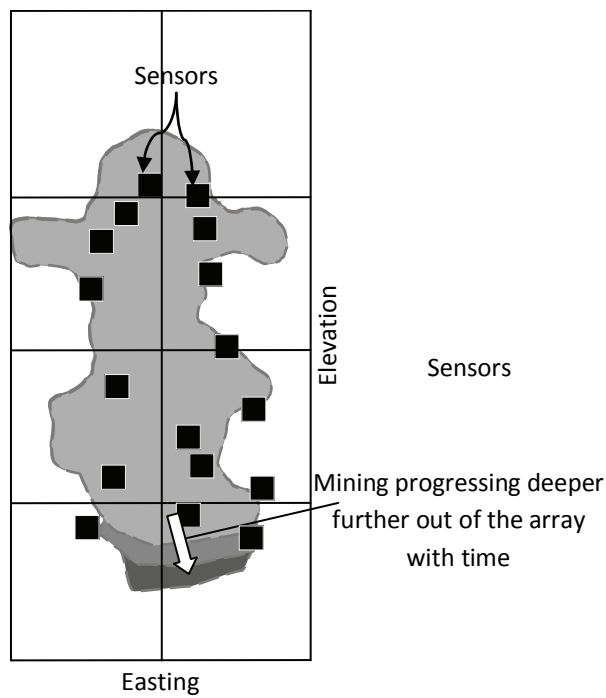


Figure 2 Illustration of mining front migrating out of the seismic array and into areas of lower system sensitivity

2 Direct evaluation of m_{min} from seismic data

The value of m_{min} can be calculated over a volume based on the calculation of the m_{min} value from the frequency magnitude distribution (FMD) for subsets of data. The data is sub-sampled according to its location. Different spatial sub-sampling schemes can be used, the simplest of which is a grid based calculation. This is the method used by Wiemer and Wyss (2000) to map the m_{min} for earthquakes in areas of Alaska, US and Japan.

A refinement of the simple grid based technique is described by Wiemer and Wyss (2000) for two-dimensional earthquake application. This scheme is not based on a fixed grid but distances between calculation nodes are determined by event density in the area. The variable grid size is necessary to achieve statistical significance of the dataset on which to base the calculation of m_{min} .

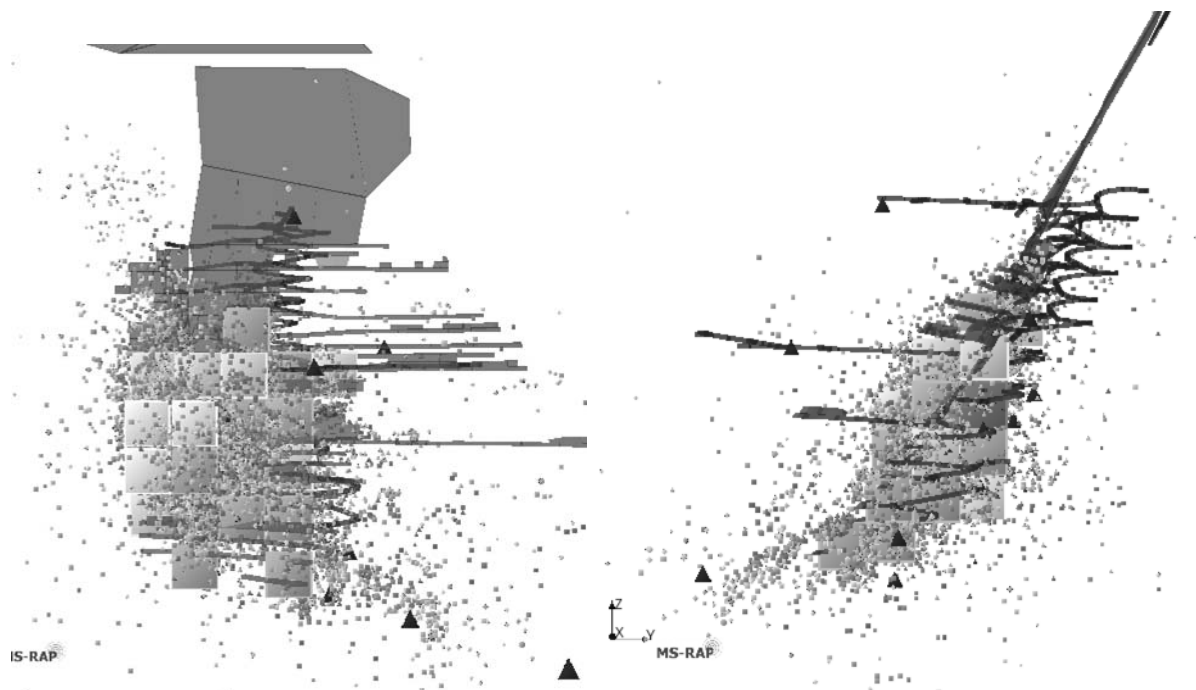


Figure 3 Results from grid based m_{min} calculation for an Australian mine

The simple fixed grid scheme was used in this work. For this, the mine volume is discretised into cubic celled grid with calculations performed on the data located in each grid cell. Figure 3 shows the recorded data in two section views for an Australian mine (referred to as Mine A). Also shown are the cells for which a value of m_{min} could be achieved. The whole volume of the mine was divided into 50 x 50 x 50 m cells, but only the cells for which reasonable values could be obtained are shown. A modest cut-off value of 50 events on which to base the m_{min} calculation were used in this case.

The strong dependence of the method on the grid size that is used is illustrated in Figure 4. The colour of the blocks represents a relative measure of the confidence in the m_{min} estimates within that block, the darker colours relates to lower numbers of data points and therefore a lower confidence in the m_{min} calculation for that particular block. It should be noted that the rock mass to the west of the mine (increasing Y-axis) is less seismically active than the east, which results in less data in this area. Although the sensitivity of the system compares well to the other areas not enough data is available to obtain values for m_{min} based on the FMD in grids in this area. The size of the cells can be increased to allow for more events to be included in each cell. This occurs, however, with a reduction in the resolution of calculation.

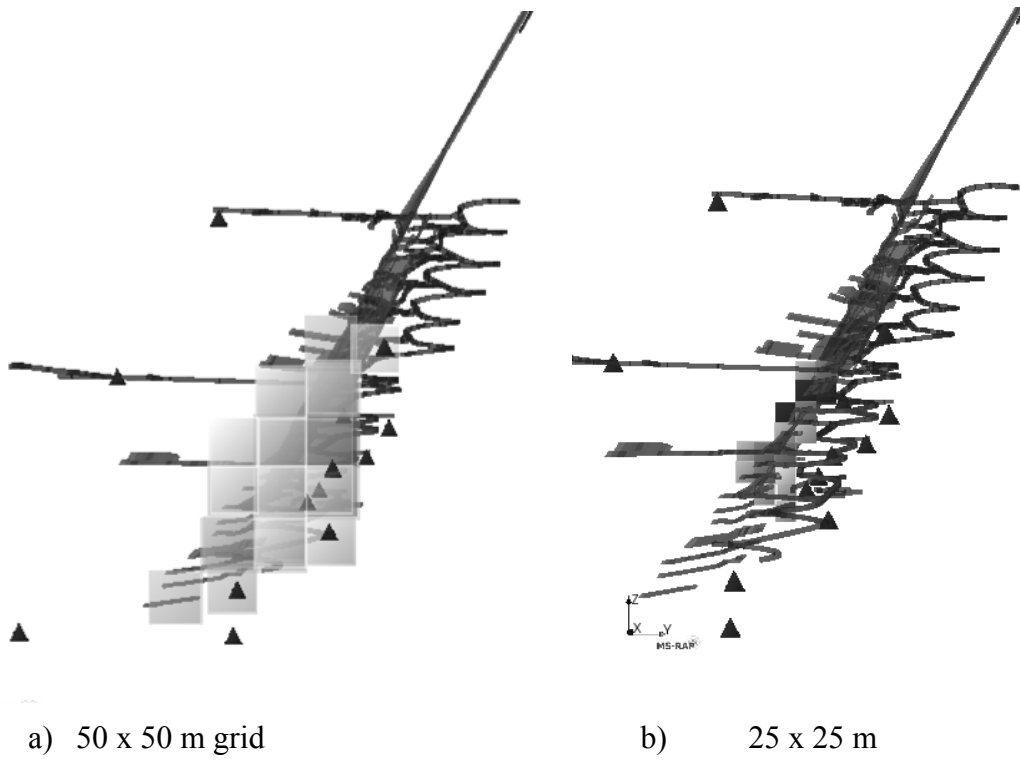
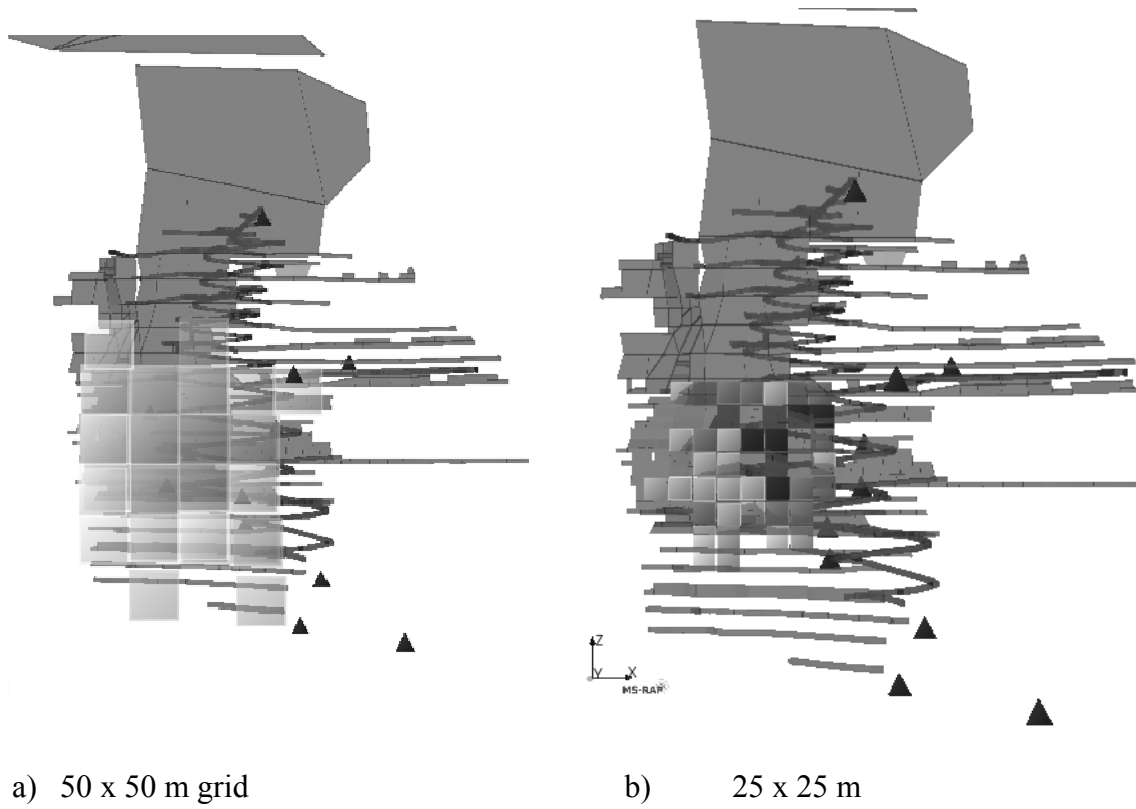


Figure 4 Results from grid based m_{min} calculation for an Australian mine

2.1 Indirect method

The major drawback of the previous method is that it is limited to the areas of high data density. This problem can be overcome by establishing a relationship between a parameter of the system geometry and the m_{min} .

Two of the most important factors determining the sensitivity of the system are the distance to the sensors and the noise level at the different sensors. As a seismic wave radiates outwards it attenuates with distance from the source. A sensor will only record the seismic wave if the amplitude of the wave is larger than the background noise level. The distance to the sensors has a first order influence on the sensitivity of the system.

A minimum of three, preferably four sensors are necessary to locate a seismic event. The minimum number of sensors to be used for locating events can be set in the system and in most Australian mines is set to four or five. Whether or not the event will be located will be determined by whether the fourth nearest or fifth depending on the setting, sensor to the source of the event is within the range at which an event of a particular size will be detected by that sensor. The distance to the fourth sensor, or fifth sensor, is therefore a good parameter related to the system geometry that can be correlated with system sensitivity. The mines from which data is used in this study used a threshold of five sensors. As a result the distance to the fifth sensor will be used here and is denoted by D5.

2.1.1 Correlation of grid based calculation of m_{min} with D5

A logical extension to the grid based assessment of m_{min} is to correlate these m_{min} values with the D5 value of each of the grid cells for which a m_{min} value were obtained. This will allow an extrapolation of the results to cells, for which m_{min} values could not be obtained.

Figure 5 shows the relationship between D5 for the grid cells shown in Figure 4 and the m_{min} value obtained from the events located in those cells. Three zones are also indicated on the plot. The lower boundary below which the system does not record any events, the upper boundary above which the system records the events and a zone which defines the area between the two boundaries with some uncertainty in the value of m_{min} .

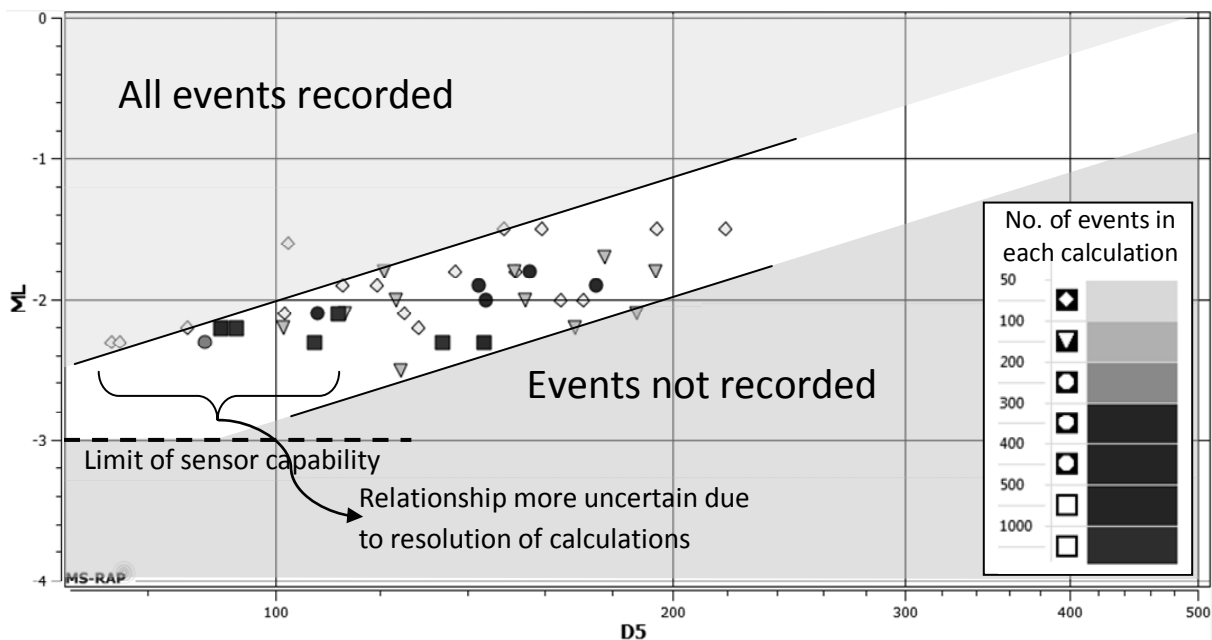


Figure 5 Relationship between D5 and m_{min} for the 50 x 50 x 50 m grid cells shown in Figure 4

Although this method allows us to obtain a relationship between the D5 and the m_{min} , it relies on a fair amount of subjective interpretation with resulting uncertainty. Each of the estimates of m_{min} in each grid is based on a limited number of data point and represents in total only a fraction of the data in the whole

database. Further uncertainty is due to the wide resolution of the calculated grids. This uncertainty is more pronounced towards the smaller values of $D5$.

This method, although allowing for the extrapolation of the m_{min} assessment to areas not covered by the direct evaluation of m_{min} , still suffers from most of the problems associated with the direct method.

2.1.2 Gridless method for obtaining a relationship between m_{min} for $D5$

Assuming a rock mass that is homogenous with respect to the seismic attenuation properties, the system sensitivity is not determined by the absolute location of the event but its location relative to the sensors. If this assumption holds, the event location in three-dimensional space becomes irrelevant and data can be analysed on the basis of its distance parameters, rather than its true three-dimensional location. Performing the m_{min} analysis on this basis, implies a superposition of data from different seismic sources with different seismic characteristics, including different values for the power law exponent, or b-value. Although important for the analysis of seismicity, the separation of seismic sources with different character is irrelevant for the analysis of the system sensitivity.

Figure 6 illustrates this concept. If the FMD of Sources 1 and 2 follow the Gutenberg–Richter (1944) relation, it follows that the combined distribution will also. The maximum value of m_{min} of the two sources will be the value of m_{min} for the combined source. Assuming a relationship exists between m_{min} and $D5$, sources at equivalent $D5$ will obtain similar values of m_{min} . Figure 6 illustrates the most-used, and in the words of Lasocki (2008), “most over-used”, Gutenberg–Richter relation. The concepts are best illustrated by assuming this relation, but are not reliant on it.

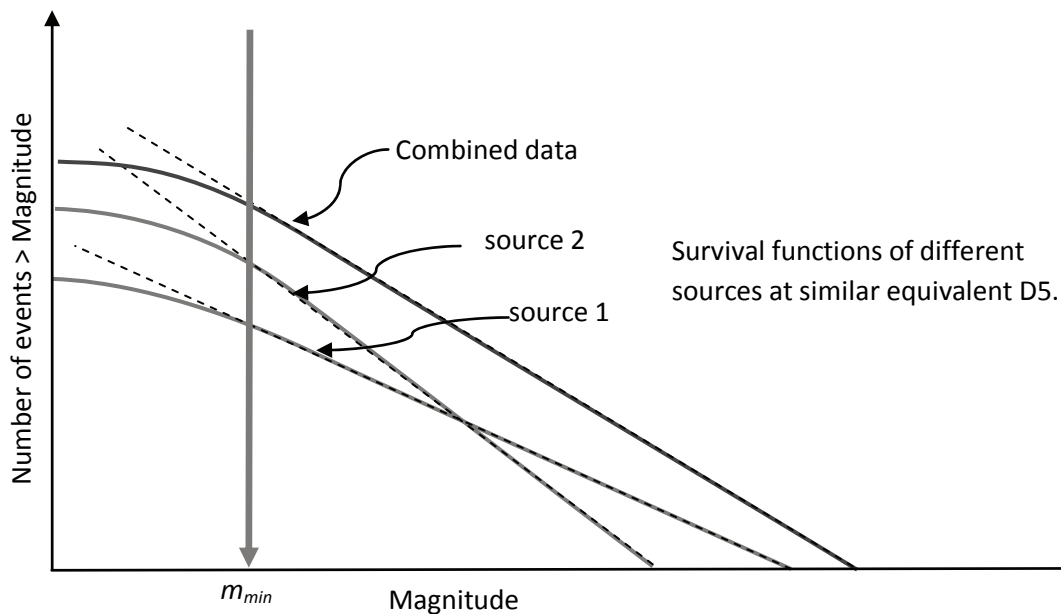


Figure 6 Illustration of the superposition of magnitude frequency data from different seismic sources at similar $D5$

Figure 7a shows a scatter plot of M_L versus $D5$ for all events recorded at Mine A. Figure 7b shows the FMD of the data with $D5$ value between 100 and 110 m. From this data, an estimate of m_{min} for the subset of data at $D5$ between 100 and 120 m can be obtained.

By performing a m_{min} calculation for subsets of data with a moving window of $D5$, the relationship between m_{min} and $D5$ is obtained as shown in Figure 8. The original data is shown as ‘x’ in the figure. Also shown as ‘+’ are the estimates of m_{min} based on the 50 x 50 x 50 m grid cells, shown in Figure 4. The shape and colour scale of the markers of the results in of the estimates of m_{min} for windows of $D5$ show the number of data points, with magnitudes greater than m_{min} in each of the $D5$ windows used in the calculations. In contrast to the grid based m_{min} estimates, the largest part of the database was used in this method resulting in a stable estimate of the m_{min} - $D5$ relationship.

A similar analysis was performed for another Australian mine (Mine B) and the resulting m_{min} -D5 relationship is shown in Figure 9.

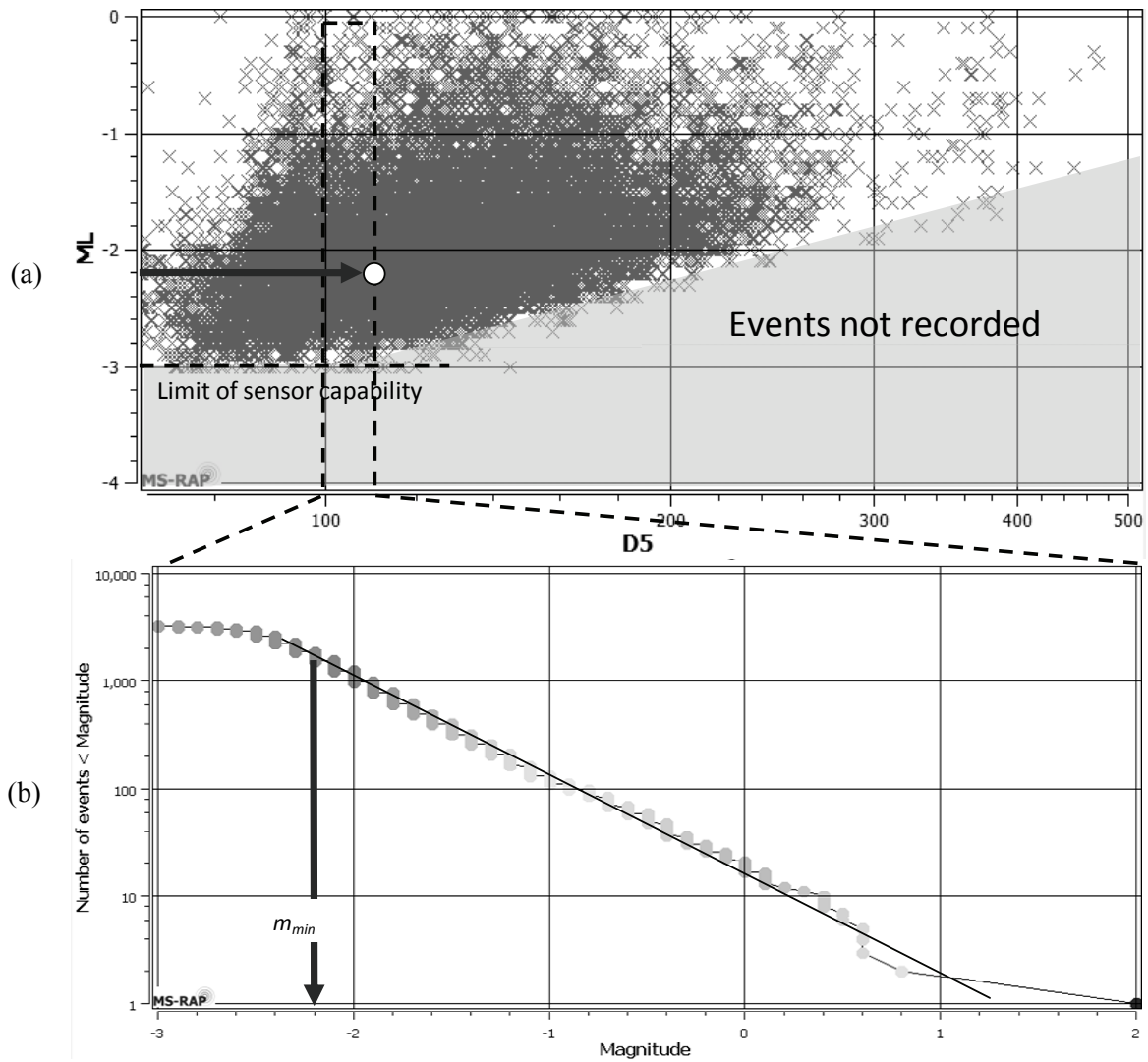


Figure 7 The relationship between D5 and m_{min} for the whole events database of Mine A (a) with calculation of m_{min} for a subsample of data with D5 between 100 and 110 m (b)

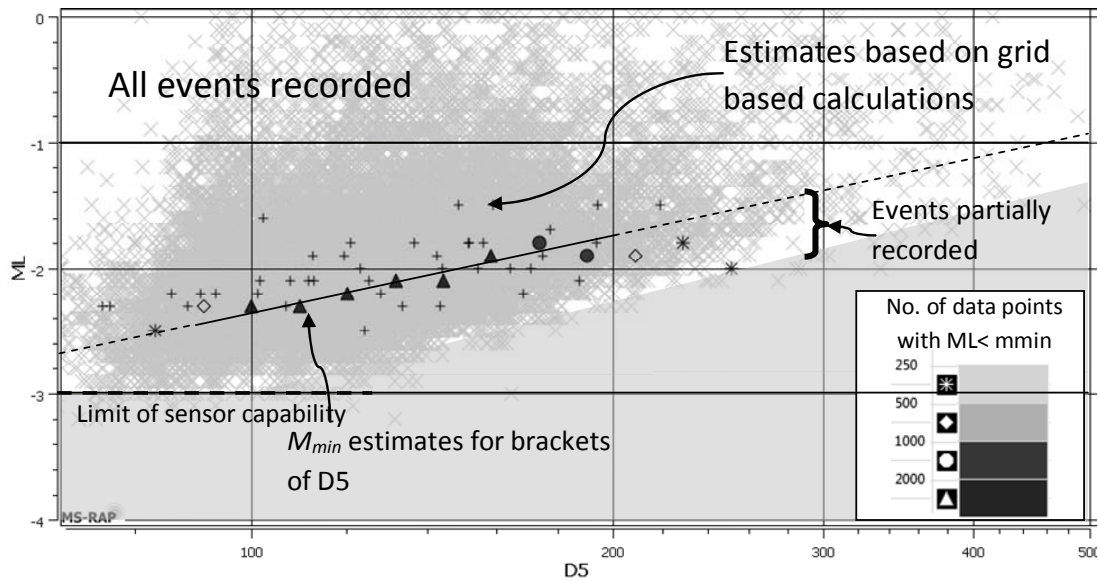


Figure 8 D5-magnitude plot for the database events of Mine A with the results of m_{min} calculations on windows of D5 data

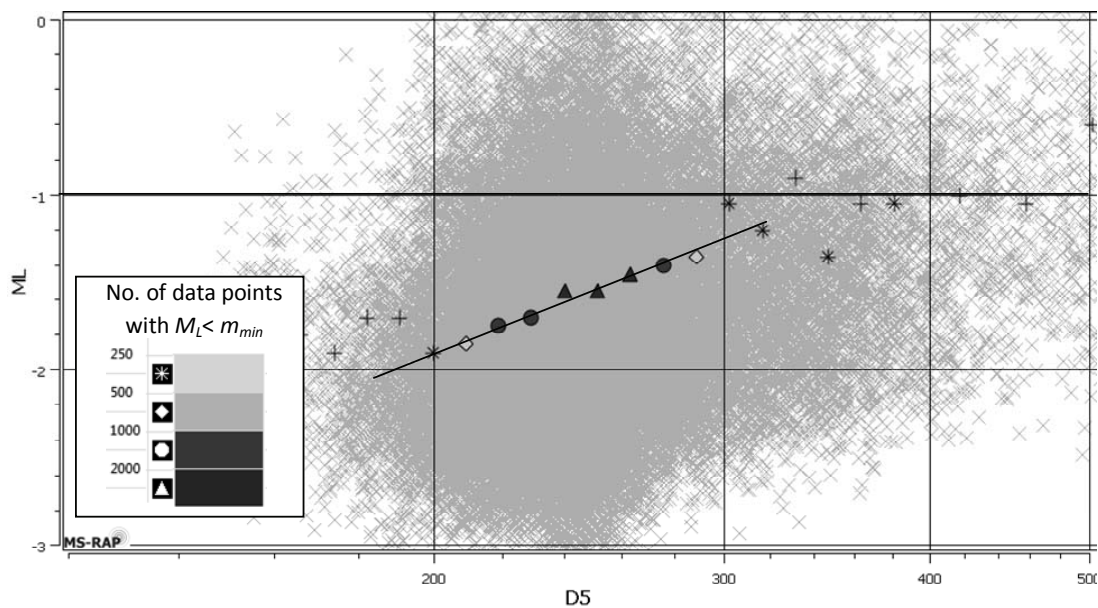


Figure 9 D5-magnitude plot for the database events of Mine B with the results of m_{min} calculations on windows of D5 data

2.2 Spatial assessment of m_{min} based on the m_{min} -D5 relationship

Having established a relationship between m_{min} and D5, this can be used to evaluate the spatial variation of m_{min} . This is done by calculating the D5 value at a point in space, and through the m_{min} -D5 relationship estimate the expected m_{min} value at that point. For illustration, the sensitivity of a seismic system at Mine B is shown on a horizontal and vertical plane through the mine (Figure 10). An extra cut-off was introduced at this point. This is the lowest magnitude that could be reliably recorded by the sensor, due to the sensor frequency response upper limit. Obtaining these cut-offs are outside of the scope of the current paper and readers are referred to Mendecki (1997). The introduction of these cut-offs is only necessary to safeguard against unrealistic low values of m_{min} where sensors are spaced very close. With the distributions of sensors used in mines, this is generally not a problem but should be included for prudence sake. The sensor density

for the example shown in Figure 10 was not dense enough to invoke this sensor lower limit. These cut-off values are shown in Figure 7 and Figure 8.

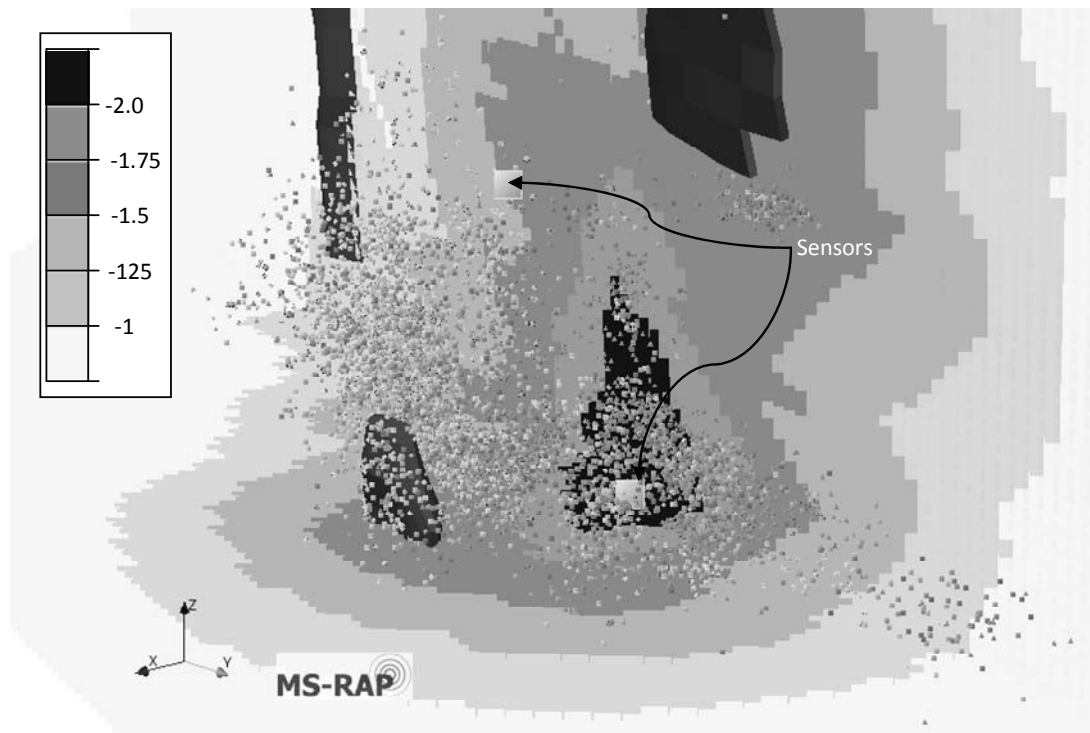


Figure 10 Contours of m_{min} at Mine B calculated from the D5 and the m_{min} -D5 relationship

D5 is a simple parameter that cannot capture the full complexity of the factors influencing the system sensitivity. Having said this, the good m_{min} -D5 correlations obtained indicate that it can be successfully used as a first order sensitivity assessment.

The D5 parameter has a physical meaning easy to understand and to visualise. It is therefore, a useful parameter for conceptualising the influence of the addition or loss of sensors on the system sensitivity in an area. With easy ‘distance-to’ calculations, the conceptual design of future system expansions can be done.

The method for obtaining a relationship between m_{min} and D5 assumes a homogenous rock mass with respect to the seismic attenuation. If this assumption is not valid, the analysis can be performed on subsets of the data, applicable to different areas of the mine. For very complex rock masses, this method may not yield good estimates and more sophisticated methods must be employed.

3 Conclusion

The direct method of m_{min} estimate for cells in a discretised volume does not appear to provide any practical solution for the assessment of system sensitivity throughout the mine volume, as this method provides only a limited number of estimates in a few discrete locations in the mine. Correlation of these m_{min} estimates with the D5 parameter provides a way of extrapolating these results for the whole mine volume. Only a small proportion of the data is used in this method for the estimates of m_{min} with resulting uncertainty in the correlation.

The gridless method for the estimation of the m_{min} -D5 relationship provides good and stable results and is based on the largest proportion of the database.

After establishing the m_{min} -D5 relationship, this can, with simple ‘distance-to’ calculations, be used to assess the system sensitivity throughout the mine. D5 is a simple parameter that cannot capture all the complex factors influencing the sensitivity of the system. In spite of its limitations, it can successfully be used for first order assessment of seismic system sensitivity and future expansion design. The D5 parameter has a physical

meaning easy to understand and to visualise and is therefore a useful parameter for conceptualising the contribution of each sensor to the overall sensitivity in a specific area.

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