Rock stresses, its controls and associated ground behaviour at the Rosebery mine, Tasmania

MF Lee  Monash University, Australia
R de Vries  MMG Limited, Australia
CF Moller  MMG Limited, Australia

Abstract
The Rosebery mine, on the west coast of Tasmania, is currently one of Australia’s deepest mines. The orebodies dip moderately (45°) and mining is currently about 1,800 m below Mount Black. As might be expected, ground conditions are challenging regarding rock stresses versus the strength of the rock mass.

Over 22 years, from 1983 to 2015, in situ rock stress measurements have been made at 16 sites at varying depths from about 500 to 1,800 m below surface. This data is documented and summarised, with depth and possible geologic and structural controls, to provide an understanding of what locally controls Rosebery’s rock stresses.

This understanding has been used to better appreciate some aspects of the mine’s challenging ground conditions.

Keywords: rock stress measurement, rock stresses, structures, ground behaviour, ground support

1  Introduction
Rock stresses are a necessary and critical input to the assessment of ground behaviour (understanding and prediction). This is particularly true for deep underground mines such as Rosebery where stresses around openings can be similar to the strength of the intact rock.

Unfortunately, the measurement of rock stresses is not always straight-forward nor cheap. Data is often only collected at ‘points’, which for practical engineering purposes must be considered to be representative of much larger volumes of rock mass. Results can be variable, especially between widely-spaced ‘point’ measurements, and, therefore, their validity and direct applicability to design issues is often questioned.

This paper documents what is presently known about rock stresses at Rosebery. Reasons for variability, between some sites, is also discussed along with the confidence that can be placed in the data, particularly as it applies to the current deep mining areas from 2,300 to 2,000 mRL (approximately 1,500–1,800 m below surface.

2  Setting and ground behaviour
Mineralisation at Rosebery was discovered in 1893. Mining commenced in about 1912, but full-scale production only started in 1936 after the Electrolytic Zinc Company of Australasia acquired the deposit. MMG Limited gained control of the operation in 2009 as part of the purchase of the Australian assets of Zinifex. The extent of past mining, and proposed future resources, beneath Mount Black are illustrated in Figure 1.
Figure 1  Perspective view of the Rosebery township and Mount Black to the north, with historical underground mining operations and drillhole traces (looking down and towards the northwest)

2.1  Geology

The Rosebery Zn-Pb-Ag massive sulphide deposit is hosted by the Mount Read Volcanics, Cambrian sediments within the Dundas Trough of western and north-western Tasmania (Martin 2004). In the immediate vicinity of the deposit, the mine sequence rocks (MS rocks) of the Mount Read Volcanics strike approximately 005°, with respect to Mine North, and dip 45° towards Mine East beneath Mount Black (Figures 2 and 3).

Note that unless specifically mentioned, all references to north are Mine North = 11.3° west of Map Grid of Australia (MGA) north.

Figure 2  Top right: typical cross-section (looking north) through the Rosebery deposit, showing multiple ore lenses and the two regional bounding faults; Mount Black Fault in the hanging wall and Rosebery Fault in the footwall. Bottom left: idealised stratigraphic column for the mine sequence rocks with details of rock units
The deposit comprises several concordant and en echelon massive sulphide lenses hosted by the transitional stratified volcanioclastics (TSV) of the MS rocks (details in Figure 2). During deposition, the TSV was irregularly intruded by porphyry sills. A thin black slate is also present in the hanging wall of the TSV. The MS rocks were subsequently variably altered, metamorphosed, foliated, folded and faulted.

The mineralisation is subparallel to bedding. Sericite alteration (Fe and Fe+Mg varieties) is common within and away from the massive sulphide contacts; nominally 5 and 20 m (stratigraphically) into the hanging wall and footwall contacts, respectively.

The MS rocks have a well-developed foliation that dips to the east, up to 15° steeper than bedding and it often has a different strike; locally ±15°. Foliation is defined by quartz, carbonates (Ca, Mg, Fe and Mn varieties) and chlorites (Fe and Fe+Mg varieties).
2.2 Specific structures

Large regionally developed bedding-parallel faults bound the MS rocks (Figure 2). Mount Black Fault, in the hanging wall, is a thick altered cataclastic breccia, but healed. Rosebery Fault, in the footwall, is a thick re-activated shear-foliation, gouge infill with clay-pug.

Other significant structures are rare at Rosebery and they are discontinuous along-strike and down-dip. Within the immediate vicinity of ore lenses, structure continuity are typically < 50 m (Figure 3). Bedding plus five cross-cutting orientation sets are loosely recognised:

1. Steep east-dipping (subparallel to foliation).
2. Steep northeast-striking shears.
4. Moderate to steep west-dipping faults.
5. Random (mostly flat-dipping joints).

Orientation Sets 1, 2 and 3 are widely spaced and are not ubiquitously developed on the scale of stoping blocks. In the vicinity of the ore lenses, and along with occasional bedding-parallel structures, their infill minerals are typically sericite plus chlorite, and associated slickensides tend to be sub-horizontal.

The west-dipping faults (Set 4) and flat-dipping joints (Set 5) are more common that the other sets, but they are also widely-spaced and often occur as en echelon arrays. Infill minerals are typically quartz plus chlorite, and associated slickensides tend to plunge steeply. West-dipping faults and moderate east-dipping structures (e.g. bedding or foliation) potentially define high ‘cathedral’ wedges in the back of strike drives. Drives must, therefore, be appropriately shaped and adequately supported.

2.3 General ground conditions — blockiness

When it has not been otherwise disturbed by mining, the Rosebery rock mass is competent for all common TSV rock types, including the porphyry sills and mineralisation. Typical rock mass quality parameters according to the Q-system (Barton et al. 1974) are given in Table 1.

Table 1 Typical rock mass parameters for common TSV rock types at Rosebery

<table>
<thead>
<tr>
<th>Host rocks and mineralisation</th>
<th>Worst</th>
<th>Average</th>
<th>Best</th>
<th>Comment (for average conditions)</th>
</tr>
</thead>
<tbody>
<tr>
<td>RQD</td>
<td>75</td>
<td>90</td>
<td>100</td>
<td>Good to excellent.</td>
</tr>
<tr>
<td>Jn</td>
<td>6</td>
<td>3</td>
<td>2</td>
<td>Foliation + discontinuous random joints.</td>
</tr>
<tr>
<td>Jr</td>
<td>1.5</td>
<td>2</td>
<td>4</td>
<td>Worst joint = foliation; smooth undulating, sometimes planar.</td>
</tr>
<tr>
<td>Ja</td>
<td>3</td>
<td>2</td>
<td>0.75</td>
<td>Worst joint = foliation; non-softening infill (quartz, carbonate, chlorite +/- sericite).</td>
</tr>
<tr>
<td>Jw</td>
<td>1</td>
<td>1</td>
<td>1</td>
<td>Drained and dry.</td>
</tr>
<tr>
<td>SRF</td>
<td>20</td>
<td>15</td>
<td>10</td>
<td>Current mining areas; high stress and moderate squeezing. Only rarely or locally influenced by a fault or weak contact.</td>
</tr>
<tr>
<td>Q'</td>
<td>6</td>
<td>30</td>
<td>267</td>
<td></td>
</tr>
<tr>
<td>Q</td>
<td>0.3</td>
<td>2.0</td>
<td>27</td>
<td></td>
</tr>
<tr>
<td>GSI</td>
<td>60</td>
<td>70</td>
<td>80</td>
<td></td>
</tr>
</tbody>
</table>

In the current mining areas, the general blockiness of the rock mass does not dictate ground behaviour.
2.4 Rock properties

Typical rock properties for Rosebery’s main TSV rock types, in the immediate vicinity of ore lenses, are given in Table 2.

Table 2 Typical rock properties for common TSV rock types at Rosebery

<table>
<thead>
<tr>
<th>Rock type</th>
<th>Density (t/m³)</th>
<th>V_p (m/s)</th>
<th>V_s (m/s)</th>
<th>E dyn (GPa)</th>
<th>E static (GPa)</th>
<th>v</th>
<th>UCS50 (MPa)</th>
<th>UTS (MPa)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Hangingwall and Footwall host rocks</td>
<td># 28</td>
<td>6</td>
<td>6</td>
<td>6</td>
<td>46</td>
<td>35</td>
<td>20</td>
<td>8</td>
</tr>
<tr>
<td></td>
<td>mean 2.80</td>
<td>5,733</td>
<td>3,179</td>
<td>73.8</td>
<td>69.1</td>
<td>0.31</td>
<td>131</td>
<td>13.0</td>
</tr>
<tr>
<td></td>
<td>std 0.09</td>
<td>627</td>
<td>170</td>
<td>8</td>
<td>20.4</td>
<td>0.08</td>
<td>34</td>
<td>4.5</td>
</tr>
<tr>
<td>Mineralised</td>
<td># 2</td>
<td>2</td>
<td>2</td>
<td>2</td>
<td>7</td>
<td>6</td>
<td>11</td>
<td>13</td>
</tr>
<tr>
<td></td>
<td>mean 4.32</td>
<td>6,183</td>
<td>2,769</td>
<td>89.6</td>
<td>122.1</td>
<td>0.26</td>
<td>158</td>
<td>7.6</td>
</tr>
<tr>
<td></td>
<td>std 0.33</td>
<td>1,091</td>
<td>171</td>
<td>1</td>
<td>49.0</td>
<td>0.08</td>
<td>44</td>
<td>2.2</td>
</tr>
<tr>
<td>Porphyry</td>
<td># 22</td>
<td>64.1</td>
<td>0.28</td>
<td>152</td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>mean 2.85</td>
<td>26</td>
<td>11</td>
<td>15</td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>std 0.10</td>
<td>9.4</td>
<td>0.07</td>
<td>61</td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
</tbody>
</table>

The prominence and pervasiveness of foliation meant that numerous tests failed by shearing, and not through intact rock, especially when foliation intersected specimens at a moderate angle of < 55. The above summary only includes tests that failed through intact rock. Uniaxial Compressive Strength (UCS) specimens had height to diameter ratios of about 2.5:1 and have been corrected to 50 mm diameter specimens using relationships developed and popularised by Hoek and Brown (1980).

Rosebery’s well-developed foliation implies that the large scale behaviour of the rock mass should be anisotropic regarding strength and probably also deformation. However, the available laboratory test data suggests that there is no significant anisotropy of the intact rock modulus (E) up to the scale of 150 mm cores.

No laboratory information is presently available regarding the shear strengths or stiffnesses of Rosebery’s structures.

2.5 Ground behaviour

In the current deep mining areas, ground conditions and support requirements for Rosebery’s ‘square-arched’ strike drives are challenging. Deformations > 300 mm are common, which have meant challenges with providing adequately supported drives (Thomas 2015; Villaescusa et al. 2014). Wall and back deformation can be high, close to advancing faces.

The key elements of ground behaviour around strike drives in highly-stressed and foliated rock masses are illustrated in Figure 4 (Beck & Sandy 2003) and summarised below:

- Shearing on moderate east-dipping structures, e.g. bedding, foliation or mineralised contacts (Figure 5).
- Induced tensile cracking away from shearing structures around the top hanging wall (exhibited as bulging) and bottom footwall corners (shown as floor heave) of the drive.
- High stress cracking through intact rock around the top footwall corner due to locally excess tangential stresses, compared to the strength of the intact rock.

Note that mechanical Dynamic (MD) bolts are used at Rosebery, not split sets as mentioned in Figure 4, and guillotining of them is an issue (Thomas 2015) due to shearing on structures and cracking around strike drives.
Similar squeezing ground has been experienced in other mines, and has been well documented by Potvin and Hadjiigeorgiou (2008). It often occurs when the following conditions are present:

- Mine openings strike subparallel (± 30°) to a prominent layering (bedded, contact, fault, foliation etc.).
- The prominent layering intersects practical ‘square’ mine openings at a low to moderate angle (< about 45°).
• The layering is weak in shear.
• Stresses are high and/or deviatoric.

To better-understand and predict (i.e. model) the severity of this style of ground behaviour on drive stability, it is important to have a good understanding of the following:

• Local stresses: principal stress magnitudes, ratios and their orientations.
• Intact rock properties for the main rock types: compressive and tensile strengths, plus deformability (E and v).
• Shear strengths and stiffnesses of the prominent layering (structures).

Despite high measured stresses in the current mining areas at Rosebery, cross-cuts and raise-bored ventilation shafts (remote from mining) do not exhibit adverse behaviour, or any other high stress behaviour, unlike the strike drives. However, discing of diamond drill core is common in sub-horizontal boreholes oriented east–west (± 45°), especially in the zone up to 15 m away from the development drives and stub cross-cuts.

3 Rock stresses

3.1 Measurements

Over a 22 year period (1983 to 2015), rock stress measurements have been made at 16 sites at varying depths from 500 to 1,800 m below surface (Figure 6). Two techniques have been used:

• Stress relief by in situ overcoring of CSIRO Hollow Inclusion (HI) Cells; 10 sites with eight successful.
• Acoustic emission (AE) method undertaken by the Western Australian School of Mines (WASM) on exploration diamond drill cores; six sites from two boreholes.

Site and measurement data, plus results, are listed in Table 3.
### Table 3: Compilation of Rosebery’s rock stress measurements

<table>
<thead>
<tr>
<th>Source</th>
<th>Site</th>
<th>Location</th>
<th>N [m]</th>
<th>E [m]</th>
<th>mRL [m]</th>
<th>Rock</th>
<th>Type</th>
<th>Technique</th>
<th>Rock properties</th>
<th>Major Principal stresses</th>
<th>Normal vertical stress (MPa)</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>BFP 1983</td>
<td>15 Level Main North Drive, 99 mRL</td>
<td>15 Level Main North Drive, 99 mRL</td>
<td>645</td>
<td>255</td>
<td>2950</td>
<td>Chloritic sericite, siliceous tuff</td>
<td>20</td>
<td>CSIRO Hi cell overcoring</td>
<td>41</td>
<td>Good</td>
<td>Magn</td>
</tr>
<tr>
<td>BFP 1999</td>
<td>NED 1120 SP</td>
<td>NED 1120 SP</td>
<td>1121</td>
<td>713</td>
<td>2450</td>
<td>Footwall volcanics (foliated)</td>
<td>26</td>
<td>CSIRO Hi cell overcoring</td>
<td>24</td>
<td>Poor</td>
<td>Magn</td>
</tr>
<tr>
<td>WASM 2006</td>
<td>W Decline DDC; R8563 24.6–27.5 m</td>
<td>W Decline DDC; R8563 24.6–27.5 m</td>
<td>1717</td>
<td>1055</td>
<td>2160</td>
<td>Quartz sericite schist</td>
<td>WASM AE</td>
<td>18</td>
<td>No rating</td>
<td>Magn</td>
<td>Dip [°]</td>
</tr>
<tr>
<td>WASM 2006</td>
<td>W Decline DDC; R8563 101.9–105.0 m</td>
<td>W Decline DDC; R8563 101.9–105.0 m</td>
<td>1710</td>
<td>1120</td>
<td>2135</td>
<td>Quartz sericite schist</td>
<td>WASM AE</td>
<td>30</td>
<td>No rating</td>
<td>Magn</td>
<td>Dip [°]</td>
</tr>
<tr>
<td>WASM 2006</td>
<td>W Decline DDC; R8563 151.0–154.0 m</td>
<td>W Decline DDC; R8563 151.0–154.0 m</td>
<td>1708</td>
<td>1180</td>
<td>2118</td>
<td>Quartz sericite schist</td>
<td>WASM AE</td>
<td>31</td>
<td>No rating</td>
<td>Magn</td>
<td>Dip [°]</td>
</tr>
<tr>
<td>WASM 2007</td>
<td>47P DDC; R8801 47.1–50.7 m</td>
<td>47P DDC; R8801 47.1–50.7 m</td>
<td>1145</td>
<td>1120</td>
<td>2136</td>
<td>Siliceous chloritic footwall host</td>
<td>WASM AE</td>
<td>36</td>
<td>No rating</td>
<td>Magn</td>
<td>Dip [°]</td>
</tr>
<tr>
<td>WASM 2007</td>
<td>47P DDC; R8801 125.1–128.8 m</td>
<td>47P DDC; R8801 125.1–128.8 m</td>
<td>1147</td>
<td>1178</td>
<td>2082</td>
<td>Siliceous chloritic footwall host</td>
<td>WASM AE</td>
<td>42</td>
<td>No rating</td>
<td>Magn</td>
<td>Dip [°]</td>
</tr>
<tr>
<td>WASM 2007</td>
<td>47P DDC; R8801 352.8–356.0 m</td>
<td>47P DDC; R8801 352.8–356.0 m</td>
<td>1151</td>
<td>1350</td>
<td>1931</td>
<td>Siliceous chloritic footwall host</td>
<td>WASM AE</td>
<td>34</td>
<td>No rating</td>
<td>Magn</td>
<td>Dip [°]</td>
</tr>
</tbody>
</table>

N, E and mRLs are with respect to Rosebery’s mine co-ordinates. Dips are positive down. Bearings are positive clockwise from Mine North = 11.3° AMG North.

Note: No useful overcore data was obtained from the two deepest 2015 attempts to measure stresses. For core diameters of 74 and 150 mm, severe discing was experienced for all sub-horizontal drilling directions. It would seem that for Rosebery’s rock mass, the practical limit of conventional rock stress measurement by overcoring has been reached.
All of the measurements were done in variably foliated and sericitised TSV rocks; sometimes pyritic. The overcoring and WASM AE measurements were done, respectively, in footwall and hanging wall rocks.

For the overcored CSIRO HI cells, the orientation of the measurement boreholes varied for each site, as did the number of cells overcored. All of the measurements were done at sufficient depths from the development openings (typically > 1.5 opening diameters), so the results would not be significantly influenced by stress-concentrations from the nearby openings.

All of the overcore testing was also done by former CSIRO Division of Geomechanics personnel who were employed at division when the HI cell was developed (late 1960s). Consequently, standard overcoring, data collection techniques and analysis procedures were used for all of the overcore measurements. Numerous QA/QC checks were routinely undertaken, redundant data was collected and standard statistical analysis and data acceptance/rejection procedures were used.

However, for a number of reasons, not all of the overcore measurements deserve the same confidence and weighting during subsequent analysis.

Table 4 outlines the experience-based qualitative criteria used to empirically rate and weight the confidence that can be placed in each site result. It is a combination of the quality and quantity of the strain and rock property data collected, which is then used to calculate each site result quoted in Table 3. In all of the following plots, the size of the symbol reflects the confidence rating of the particular site rock stress measurement.

### Table 4  Experience-based qualitative criteria used to rate and weight the confidence of rock stress measurements, by overcoreing CSIRO HI Cells (from Lee 1997)

<table>
<thead>
<tr>
<th>Rating</th>
<th>Weighting</th>
<th>Comments</th>
</tr>
</thead>
</table>
| **Excellent** | 8 | • At least three good individual overcores of CSIRO HI cells (≥ 33 strains), good rock property data for each, and no experimental difficulties.  
• Variable results between adjacent cells are not considered to be an issue, rather, a characteristic of the local rock mass.  
• Good rock property data is required, and ideally not significantly different between adjacent cells.  
• Data collection from two or more boreholes is considered to be better than from one borehole. |
| **Good** | 6 | • At least two good individual overcores (24 to 32 strains).  
• A third cell may have had gluing and/or electrical problems.  
• Significantly different rock properties between some cells, and/or anisotropic or micro-cracked rock.  
• It is difficult for stress measurements using the borehole slotting or individual tests by hydraulic fracturing to be rated any higher than good. |
| **Fair** | 3 | • Only two overcores and incomplete due to some gluing and/or electrical problems (16 to 23 strains).  
• Technical difficulties with the overcoring, e.g. groundwater, glue bonding, discing problems, transient temperature effects, but with good rock property data.  
• Alternatively, poor rock property data and up to 24 good overcore strains. |
| **Poor** | 1 | • Only one overcore or poor data from two overcores and associated biaxial tests (< 15 strains).  
• High variability between individual tests. Poor glue bonding. |
Although the WASM AE data has been shown in plots (open symbols in Figures 7, 8 and 9), this data has been ignored when developing stress relationships for Rosebery. Their orientations and magnitudes suggest slightly different relationships and ratios between principal stresses, compared to the overcore data.

![Principal Stress Orientations](image)

**Figure 7** Principal stress orientations, for CSIRO HI + WASM AE data at Rosebery

![Principal Stress Magnitude](image)

**Figure 8** Principal stress magnitudes versus normalised ‘depth’ (3,500 – mRL), for CSIRO HI + WASM AE data at Rosebery
Managing challenging conditions

Figure 9  Principal stress magnitudes versus the sum of the principal stress magnitudes, for CSIRO HI + WASM AE data at Rosebery

3.2  Relationships

3.2.1  Principal stress orientations versus Rosebery’s foliation

Rosebery’s measured principal stress orientations are plotted in Figure 7. Site ratings (weightings) were used to contour the data.

The minor principal stress \(\sigma_3\) invariably plunges steeply to the east (60/095), subparallel to the dip-direction of Rosebery’s foliation.

As their measured magnitudes are often similar, the orientations of Rosebery’s major and intermediate principal stresses \(\sigma_1\) and \(\sigma_2\) tend to swap between plunging shallowly west (30/262), sub-perpendicular to foliation, and shallowly north (06/356), subparallel to the strike of foliation.

3.2.2  Principal stress magnitudes versus ‘depth’

As the mine plunges beneath Mount Black (see Figure 1), various schemes were investigated to try to ‘correct’ the measured stresses for topographical effects. Unfortunately, they were all inconclusive and unsatisfactory. The most useful depth versus stress relationships (Figure 8), especially for the current deep mining area, was achieved by simply plotting principal stress magnitudes versus a normalised mine RL, i.e. 3,500 mRL.

Although poorly-constrained, simple linear principal stress magnitude versus ‘depth’ relationships are suggested, it must be remembered that the measured intermediate and major principal stresses have a
tendency to swap orientations. Care is, therefore, necessary when applying these relationships to assess engineering designs.

The shallow westerly dip of the major principal stress and its high magnitude in the current mining areas, compared to the strength of the rock mass, imply excessive stresses and cracking through intact rock around the top footwall corner of Rosebery’s strike drives, as is commonly observed. However, stress concentrations in the walls of cross-cuts and raises are not yet high enough to initiate cracking, as is also observed.

3.2.3 Principal stress magnitude ratios

Despite extensive validation of in situ rock stress measurements by overcoring, at numerous operations against observed ground behaviour, doubt is invariably voiced about their validity when results do not neatly fit theoretical or text-book ideas about rock stresses.

One method of checking the consistency of a set of rock stress measurement data, from a similar geotechnical domain, is to plot principal stress magnitudes versus their sum (σ₁ + σ₂ + σ₃). This is a normalisation process. It often gives well-constrained linear relationships between the principal stresses, as is suggested in Figure 9, irrespective of their orientations or variations with depth and along strike.

Harrison et al. (2007) noted that these ratios are not a mathematical obscurity of tensors and that slightly different principal stress ratios occur in different geological (i.e. geotechnical) environments.

3.2.4 Stresses versus structures

Given the active loading of tectonic plates and the fixed structural character of the Rosebery rock mass, it is argued that there are only certain maximum principal stress ratios that the rock mass can sustain. The ‘dynamic’ process involves shearing on structures in geologic-time, i.e. re-activation of structures (as aseismic creep, or perhaps earthquakes), redistribution of loads and the modification of local stresses to the maximum principal stress ratios that the local rock mass can sustain. Those ratios are very likely to be a function of the orientation and character (shear strengths and stiffnesses) of the local structures.

‘Structures’ are often cited as the reason why measured stresses are different to expected well-behaved text-book trends. It is argued in the following section that this is true at Rosebery.

3.2.4.1 Rosebery Fault

For each of the measured stresses, pre-mining normal and shear stresses acting on Rosebery Fault are plotted in Figure 10. Note that to calculate pre-mining normal and shear stresses, which were presumably at-least at limiting equilibrium with external loads prior to mining (gravitational, tectonic etc.), water pressures corrections are required. Mining effectively drains and depressurises the rock mass, thereby making it a little stronger and less prone to shearing on structures.

![Figure 10 Pre-mining normal versus shear stresses, per overcore rock stress measurement at Rosebery Fault](image-url)
The upper limit of the data suggests a low long-term (geologic) shear strength for Rosebery Fault (cohesion, \(c = 4\) MPa and friction angle, \(\phi = 10^\circ\)), which, given its weak clay-pug infilled nature, is not a surprise.

The WNW-ESE tectonic loading (with respect to true north) of the Australian Tectonic Plate through Tasmania (Hillis & Reynolds 2003) acts at about 45° to Rosebery Fault’s strike and dip. There should, therefore, be no surprise about the importance and control of Rosebery Fault, and other similar regionally-developed structures, on the style of the mine’s stress field. The local stresses must be sub-aligned with Rosebery Fault.

Measurements at three of the overcore sites have ‘high’ combinations of pre-mining normal and shear stresses; i.e. the shallow 15 Level and 19 ‘F1’ Sublevel sites, and the deep 47K site. It’s therefore also no surprise that occasional seismicity has been monitored near Rosebery Fault opposite the 47K measurement site in response to mining in the adjacent K, N and P Lens.

### 3.2.4.2 Bedding and structure orientation sets 1, 2 and 3

Pre-mining normal versus shear stresses acting on bedding and structure orientation sets 1, 2 and 3 are shown in Figure 11. The upper limit to the data also suggests a low long-term (geologic) shear strength for these structures (\(c = 9\) MPa and \(\phi = 10^\circ\)), which typically have sericite plus chlorite infill and sub-horizontal slickensides.

![Figure 11 Bedding and structure orientation sets 1, 2 and 3: top, pre-mining shear versus normal stresses on structures, per stress measurement; bottom left, structure orientations for structures that have implied cohesions > 5 MPa; and, bottom right, calculated shear directions for those structures](image)

Figure 11 Bedding and structure orientation sets 1, 2 and 3: top, pre-mining shear versus normal stresses on structures, per stress measurement; bottom left, structure orientations for structures that have implied cohesions > 5 MPa; and, bottom right, calculated shear directions for those structures

The orientations of the few structures that have high implied cohesions (> 5 MPa) have been plotted stereographically and are also shown in Figure 11. The calculated directions of shearing on these structures, as suggested by the nearby stress measurements, have also been shown stereographically as trends and plunges. Interestingly, they are all sub-horizontal, like the mapped slickensides for bedding and orientation sets 1, 2 and 3.
3.2.4.3 Steep west-dippers and random

Pre-mining normal versus shear stresses acting on steep west-dipping faults and random structures are shown in Figure 12. The upper limit to the data suggests a higher long-term (geologic) shear strength (c = 8 MPa and φ = 19°) for these structures, which typically have stronger quartz plus chlorite infill and steeply-plunging slickensides.

![Graph showing shear stress vs normal stress](image)

**Figure 12** Steep west-dipping faults and random structures: top, pre-mining shear versus normal stresses on structures, per stress measurement; bottom left, structure orientations for structures that have implied cohesions > 5 MPa; and, bottom right, calculated shear directions for those structures.

The orientations of the few structures that have high implied cohesions (> 5 MPa) have been stereographically plotted and shown in Figure 12. They all dip moderate to steeply west. The calculated directions of shearing on these structures, according to nearby overcore rock stress measurements, have also been shown stereographically as trends and plunges. Most directions are sub-horizontal, but some have shallow to moderate plunges similar to what has been observed on these structures underground.

3.2.4.4 Associated seismicity

The 47K and 47W decline measurements were conducted at similar RLs, respectively, in the footwall of K and W Lenses (Figure 6), but very different stresses were measured at these sites (Table 4). As shown in Figures 11 and 12, just a few specific structures near both sites have high shear stresses. Mining-induced seismicity has been noted in both areas due to high stress cracking/failure of intact rock around drives and stope abutments, but also due to shearing on a few specific structures, as an inspection of stress versus structure relationships would suggest. Early (relative to stoping) and wide-spread shearing on specific footwall structures was a ‘feature’ of W Lens (Figure 13), where much higher shear stresses are predicted on structures.
4 Conclusion

Over the past 22 years, the Rosebery mine has progressively invested in the collection of rock stress measurement data to assist with understanding key elements of ground behaviour and as input to stope and regional stability assessments. As often happens at some mines, the data is variable with depth and along strike, and it does not conform to simple ‘text-book’ ideas of what rock stress trends should be. It has, therefore, been tempting to dismiss and ignore some of the measurements, by citing any number of unsubstantiated reasons.

The Rosebery rock stress measurements, by overcoring CSIRO HI cells, have been reviewed and empirically rated as to the confidence that can be placed in each measurement (Table 4). This data was then used to investigate and make the following general conclusions about principal stress relationships at Rosebery, with depth and interactions with structures.

- Depth versus principal stress orientation and magnitude relationships for the current (deep) mining areas are (normalised with respect to Mine mRL):

  - $\sigma_1 = 6 + 0.055 (3,500 - \text{mRL})$ MPa; plunging shallowly west (30/262).
  - $\sigma_2 = 3 + 0.038 (3,500 - \text{mRL})$ MPa; plunging shallowly north (06/356).
  - $\sigma_3 = 0.026 (3,500 - \text{mRL})$ MPa; plunging steeply to the east (60/095).

  The principal stress orientations are aligned with Rosebery’s ubiquitous foliation (60/090). The $\sigma_1$ and $\sigma_2$ orientations often swap, probably in response to the local presence, or absence, of specific cross-cutting structures. A suitable method of accounting for the effect of Mount Black was not found, but at the current mining depths its influence is not considered to be significant.

- Unlike the depth relationships, ratios between the principal stress magnitudes are well-constrained:

  - $\sigma_1 / \sigma_2 = \sigma_2 / \sigma_3 = 1.5$.
  - $\sigma_1 / \sigma_3 = 2.2$.
• The footwall Rosebery Fault (thick with clay-pug infilled), in association with perhaps similar regionally-developed faults, is likely to be the main controlling influence on the style of rock stresses at Rosebery. The data suggests a long-term (geologic) shear strength for the fault of $c = 4$ MPa, $\phi = 10^\circ$.

• Locally specific and discontinuous cross-cutting structures have a secondary control on Rosebery’s stresses. Their specific presence/absence and orientations are likely to dictate local principal stress orientations, magnitudes and ratios. The data implies the following long-term (geologic) shear strengths for these structures, for different infill materials:
  
  ○ Bedding and orientation sets 1, 2 and 3 (sericite plus chlorite); $c = 9$ MPa, $\phi = 10^\circ$.
  ○ Steep west-dipping faults and random structures (quartz plus chlorite); $c = 8$ MPa, $\phi = 19^\circ$.

Mapped slickensides on these structures favourably correlate with shear directions predicted by nearby rock stress measurements.

• Recent mining-induced seismicity (style and position) in the footwall of the current mining blocks (K and W Lenses), correlates well with specific structures that have high shear stresses.

• The above stress relationships explain the challenging squeezing ground behaviour experienced in Rosebery’s strike drives, and the lack of adverse ground behaviour in cross-cuts and raises.

• The implied long-term shear strengths of Rosebery’s structures should be used to further-investigate the chance of shearing on particular structures around development and stoping blocks.

The WASM AE results indicated that there was slightly different relationships and different ratios between the principal stresses compared to the overcore data. The WASM AE results were not used to assess relationships.

Of course, similar reviews and analyses of rock stress measurement data, versus geology, at other mines is likely to lead to similar useful insights into local controls on rock stresses and ground behaviour.

**Acknowledgement**

The authors wish to express their appreciation to MMG Limited for permission to publish and present this paper.

**References**


